

Time dependent seismicity along the western coast of Canada

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Abstract

Decelerating generation of intermediate magnitude earthquakes (preshocks) in a narrow region (seismogenic region) and accelerating generation of relatively larger such earthquakes in a broader region (critical region) has been proposed as an appropriate model for intermediate-term earthquake prediction. We examined the seismic activity which preceded the $M_w=7.7$ (October 28, 2012) thrust event that occurred off the west coast of Haida Gwaii, Canada (formerly the Queen Charlotte islands), by applying the decelerating-accelerating seismic strain model. We found that this mainshock was preceded by a pronounced accelerating seismic sequence with the time to the mainshock, as well as by an equally easily identifiable decelerating seismic sequence. Both precursory seismic sequences occurred in different space, time and magnitude windows. The behavior of previous mainshocks that occurred close to the 2012 earthquake was also examined by the time and magnitude predictable regional model.

An attempt was also made to identify such seismic strain patterns, which may also be related to the generation of strong mainshocks along the western coast of Canada.

Introduction

Strong and large earthquakes along the western coast of Canada are not uncommon. During the instrumental period (since 1898, when the first seismograph of Milne type was installed at Victoria) several $M \geq 7.0$ events occurred there. The majority of the strong earthquakes in this area are associated with the motion between the three major lithospheric plates, namely Pacific, North America and Juan de Fuca (Figure 1). The boundary between Pacific and North America plates is a right lateral transform fault (Queen Charlotte Fault, QCF) that extends from Vancouver Island up to Alaska and the Fairweather Fault (FF), whereas the boundary between Juan de

Fuca and North America plates corresponds to the Cascadia subduction zone (CSZ), where the giant M9 earthquake occurred in 1700.

The present study was motivated by the occurrence of the 2012 Haida Gwaii islands (formerly the Queen Charlotte islands) earthquake ($M=7.7$). This was the first major thrust event recorded along the strike-slip QCF.¹ In this article we present the results of the application of two time-dependent seismicity models in an attempt to retrospectively predict the 2012 mainshock. Moreover, we apply both models searching for future strong earthquakes along the western coast of Canada.

Models applied and data

The first of the two time-dependent seismicity models is based on the triggering of a mainshock by its preshocks and is called decelerating-accelerating seismic strain (D-AS) model. There is reliable evidence that an increase in the occurrence rate of intermediate-magnitude shocks has been observed in a broad area before strong earthquakes.²⁻⁸ Recently, however, skepticism has been expressed on the reliability of these observations,^{9,10} which is discussed later. Quantification of the accelerating pattern of these earthquakes that occur in this broad region before a mainshock,¹¹ showed that the cumulative Benioff strain, $S(t)$, can be expressed by the following power law:

$$S(t) = A + B(t_c - t)^m \quad (1)$$

where t_c is the origin time of the mainshock and A , B , m , are parameters calculated by the available data (with $m < 1$, $B < 0$). The quantity $S(t)$, which is considered as a measure of the preshock seismic deformation at time t , is

defined as $S(t) = \sum_{i=1}^{n(t)} E_i^{1/2}$, where E_i is the

seismic energy of the i th preshocks and $n(t)$ is the number of preshocks occurred up to time t . On the other hand, it has also been observed that in the narrow (focal) region of an ensuing mainshock, a seismic excitation is followed by a drop of seismicity, *i.e.*, a seismic quiescence period.^{12,13} Global data were used¹⁴ to show that intermediate magnitude preshocks in the focal region form a decelerating pattern and that the time variation of the cumulative Benioff strain up to the mainshock also follows a power-law (relation 1) but with a power value larger than one ($m > 1$). That is, this pattern of decelerating strain in the focal region is formed of a seismic excitation followed by a decrease of seismicity of intermediate magnitude preshocks.

The formulation of the D-AS model for intermediate-term earthquake prediction was based on the examination of the patterns described above, which preceded strong

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($M=6.3-9.0$) shallow ($h < 100$ km) mainshocks worldwide.¹⁵ Tests performed on synthetic catalogues^{15,16} and retrospective predictions of recent strong mainshocks have been used to evaluate the model whereas forward tests led to the successful intermediate-term prediction of two strong earthquakes in the Aegean.^{17,18} During the formulation of the D-AS model it was observed that each of the investigated mainshocks was preceded by a decelerating preshock sequence, generated in a relatively small region (seismogenic region) where the mainshock is also located, and by an accelerating preshock sequence generated in a broader region (critical region) and that both precursory seismic sequences have predictive properties, related to the ensuing mainshock. Decelerating preshocks occur in different time, magnitude and space windows than the accelerating preshocks. The latter start earlier than the former ($t_{sa} > t_{sd}$), and their magnitudes are larger the magnitudes of the decelerating

preshocks. The strain acceleration, q_a , as well as the strain deceleration, q_d , (also called quality indexes), vary with the time to the mainshock.¹⁷ Both start with low values, attain their largest values several years before the mainshock occurrence and cease gradually about three years before the mainshock, *i.e.*, the seismic activity declines in the critical region and increases in the seismogenic region.

The radii r (km) of the seismogenic region and R (km) of the critical region, which are assumed circular, are given by the relations:¹⁵

$$\log r = 0.23 \cdot M - 0.14 \cdot \log s_d + 1.40 \quad \sigma = 0.15 \quad (2)$$

$$\log R = 0.42 \cdot M - 0.30 \cdot \log s_a + 1.25 \quad \sigma = 0.16 \quad (3)$$

with M being the mainshock magnitude and s_d and s_a (in $J^{1/2}/y \times 10^4 \text{ km}^2$) is the Benioff strain in the seismogenic and critical region, respectively.

The following two relations hold also for the decelerating and accelerating preshock sequences:¹⁵

$$\log(t_c - t_{sd}) = 2.95 - 0.31 \cdot \log s_d \quad \sigma = 0.12 \quad (4)$$

$$\log(t_c - t_{sa}) = 4.60 - 0.57 \cdot \log s_a \quad \sigma = 0.17 \quad (5)$$

where t_{sd} and t_{sa} are the start times (in years) of the decelerating and accelerating preshock sequence, respectively.

The curvature parameter, C , has been proposed¹⁹ as a measure of accelerating strain release and equals to the ratio of the RMS error of the power-law fit (Eq. 1) to the corresponding linear fit error. C takes positive values smaller than 1, becomes equal to 1 for linear fit and decreases when the accelerating Benioff strain release becomes more intense. A quality index, q_a , has been defined to measure the intensity of the accelerating pattern:²⁰

$$q_a = \frac{P_a}{m_a \cdot C_a} \quad (6)$$

where m_a is the exponent in relation (1), C_a is the curvature parameter and P_a is the probability that an accelerating pattern fulfills relations (3) and (5). A similar relation quantifies the intensity, q_d , of the decelerating strain:

$$q_d = \frac{P_d \cdot m_d}{C_d} \quad (7)$$

with m_d being the exponent in relation (1) for the decelerating pattern, C_d is the curvature

parameter and P_d is the probability that a decelerating strain release fulfills relations (2) and (4). Both quality indexes q_d and q_a are very useful in searching for decelerating and accelerating seismicity patterns since they attain their largest values at the seismogenic and critical region, respectively. Global observations¹⁵ resulted in the following cut-off values of the parameters that describe the decelerating and the accelerating precursory seismic sequences:

$$C_d \leq 0.60, \quad P_d \geq 0.45, \quad q_d \geq 3.00, \quad 2.5 \leq m \leq 3.5 \quad (8)$$

$$C_a \leq 0.70, \quad P_a \geq 0.45, \quad q_a \geq 3.00, \quad 0.25 \leq m \leq 0.35 \quad (9)$$

The second of the models applied in the present work makes use of interevent times of strong earthquakes in an area, which depend on the rate of tectonic loading. However, since such earthquakes, that occur on a single fault, usually have recurrence times much larger

than the relatively short span of the earthquake catalogues used, it is preferable to consider seismogenic sources, *i.e.*, circular regions that include, in addition to the main fault where the largest mainshock occurs, other smaller faults where smaller mainshocks occur. On the basis of this idea, the time and magnitude predictable regional (TIMAPR) model has been proposed,²¹ which makes use of numerous interevent times of strong earthquakes (mainshocks) generated in a region, in order to investigate the time-dependent seismicity of this region. This model has been developed by considering a large sample of global data,²¹ which was used to derive the following two equations that relate the interevent time, T_i (in years), until the next mainshock and its expected magnitude, M_n to the magnitude, M_p , of the previous mainshock in the region examined, the long-term seismicity level, S_d (the seismic strain rate in $Joule^{1/2}$ per year) and the minimum mainshock magni-

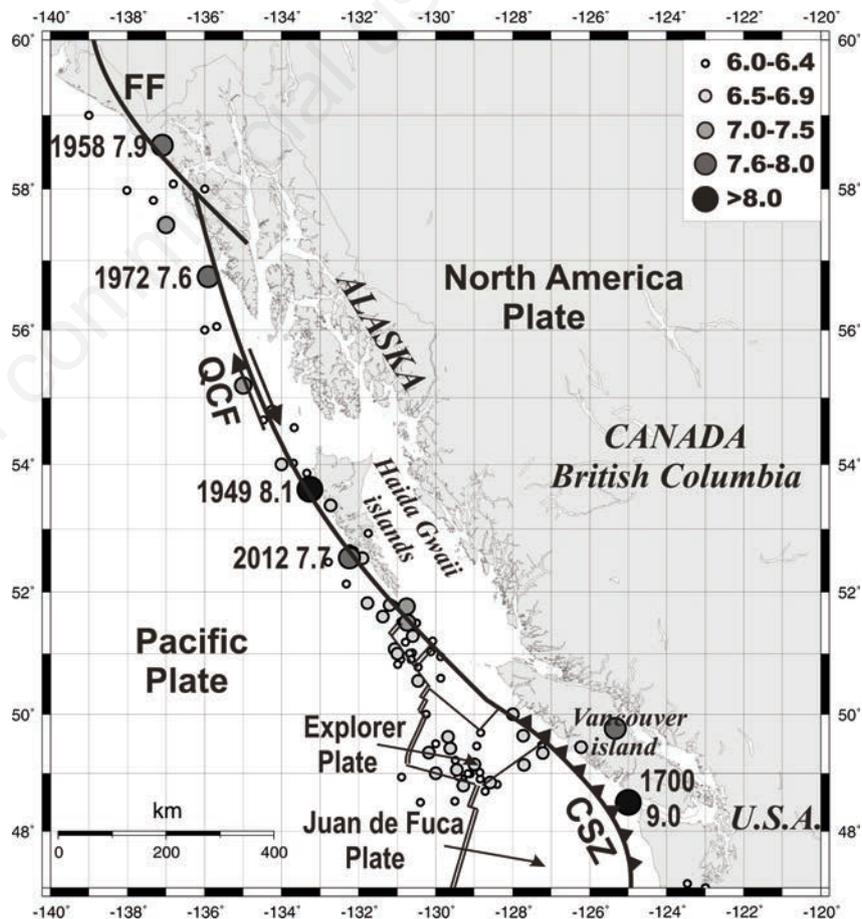


Figure 1. Main tectonic features of the studied area. FF, Fairweather Fault; QCF, Queen Charlotte Fault; CSZ, Cascadia Subduction Zone. Epicenters of all earthquakes with $M \geq 6.0$ that occurred after 1912 are shown along with the epicenter of the giant M9 of 1700 in Cascadia.

tude, M_{\min} , considered:

$$\log T_i = 0.19 \cdot M_{\min} + 0.33 \cdot M_p - 0.54 \cdot \log S_d + q \quad (10)$$

$$M_f = 0.73 \cdot M_{\min} - 0.28 \cdot M_p + 0.46 \cdot \log S_d + w \quad (11)$$

The mean values of q and w and their standard deviations σ_q and σ_w can be calculated by the available data for a certain seismogenic source. It was showed²² that the ratio σ/T_i of the observed interevent time, T_i , to the calculated, T_i (by Eq. 10), follows a lognormal distribution with a mean equal to zero. The application of this model requires a declustered catalogue of mainshocks, which is derived after adopting an appropriate time window, defined by the ratio σ/T (where T is the mean interevent time and s its standard deviation), since this ratio is a measure of seismic clustering. For ratio values smaller than 0.5 an earthquake catalogue exhibits quasi-periodic behaviour.²³ Global data show that for a time window $\Delta t \geq 15$ years this ratio becomes smaller than 0.5²⁴

The procedure followed for creating a mainshock catalogue in a seismogenic source through declustering starts after considering the largest earthquake of the available complete earthquake catalogue of the source as the first mainshock. This mainshock and its associated shocks (earthquakes of the original catalogue that occurred in the seismogenic source within a time window ± 15 years from the mainshock origin time) are excluded from the original catalogue. Then, the largest earth-

quake of the remaining catalogue is considered as a mainshock and its associated shocks, defined as previously, are also excluded from the catalogue. This continues till no event with magnitude larger than a certain cut-off magnitude M_{\min} remains in the original catalogue. For the identification of a circular seismogenic source the examined area is covered by a dense grid of geographic points. Each point is considered as the center, K , of a circular seismogenic source and circles with varying radius (*e.g.*, $r=30$ -200 km) are defined. The earthquake catalogue for each circle is declustered according to the procedure described previously and relations (10) and (11) is applied on the resulting mainshocks. The optimum seismogenic source is selected for the circle (K , r)

which the optimization factor, $OP = \frac{\sqrt{N}}{\sigma_q}$,

where N is the number of interevent times, has the largest value ($N \geq 3$). In the present case the geographic point with the highest OP value in the vicinity of the 2012 earthquake is 51.6°N - 131.0°W , with $OP=19.35$ and $r=180$ km. The declustered mainshock catalogue within this seismogenic source, which will be subsequently used for the retrospective prediction of the 2012 event, includes the following shocks: 1912 $M=6.5$, 1929 $M=7.0$, 1949 $M=8.1$, 1970 $M=7.4$, 1992 $M=6.7$.

Data used in the present work have been extracted from the Seismic Hazard Earthquake Epicenter File (SHEEF) for the period 1627-1991²⁵ and from the National Resources

Canada on-line bulletin (<http://earthquakes-canada.nrcan.gc.ca/stdon/NEDB-BNDS/bull-eng.php>, last accessed: July 23, 2014) for the period 1992-2014. The studied area is bounded by the 120°W - 140°W meridians and 47°N - 57°N parallels. Earthquake magnitudes are mainly expressed in the M_L , M_S and m_b scales and M_w for recent events. Although earthquake detectability studies²⁶ suggest that the earthquake catalogue for the examined area may be assumed to be complete for $M \geq 5.0$ since 1951 and $M \geq 4.0$ since 1971, we used the ZMAP software²⁷ as well as the traditional frequency-magnitude distribution²⁸ to check space and time variations of magnitude completeness. We found that the data are complete for the whole area under investigation for $M \geq 4.8$ for the time interval 1960-1990 and for $M \geq 4.0$ for the time interval 1991-2014. To ensure homogeneity of the catalogue in respect to the magnitude, m_b and M_S magnitudes were converted into the moment magnitude scale, M_w , by appropriate relations derived by the use of global data,^{29,30} whereas M_L magnitudes were converted to M_w using the relations valid for western Canada.³¹

Results

Previous studies^{15,16,24,32} showed that the centers $F(\varphi, \lambda)$ and $Q(\varphi, \lambda)$ of the seismogenic and critical regions, respectively, are located in areas bounded by the two parallels

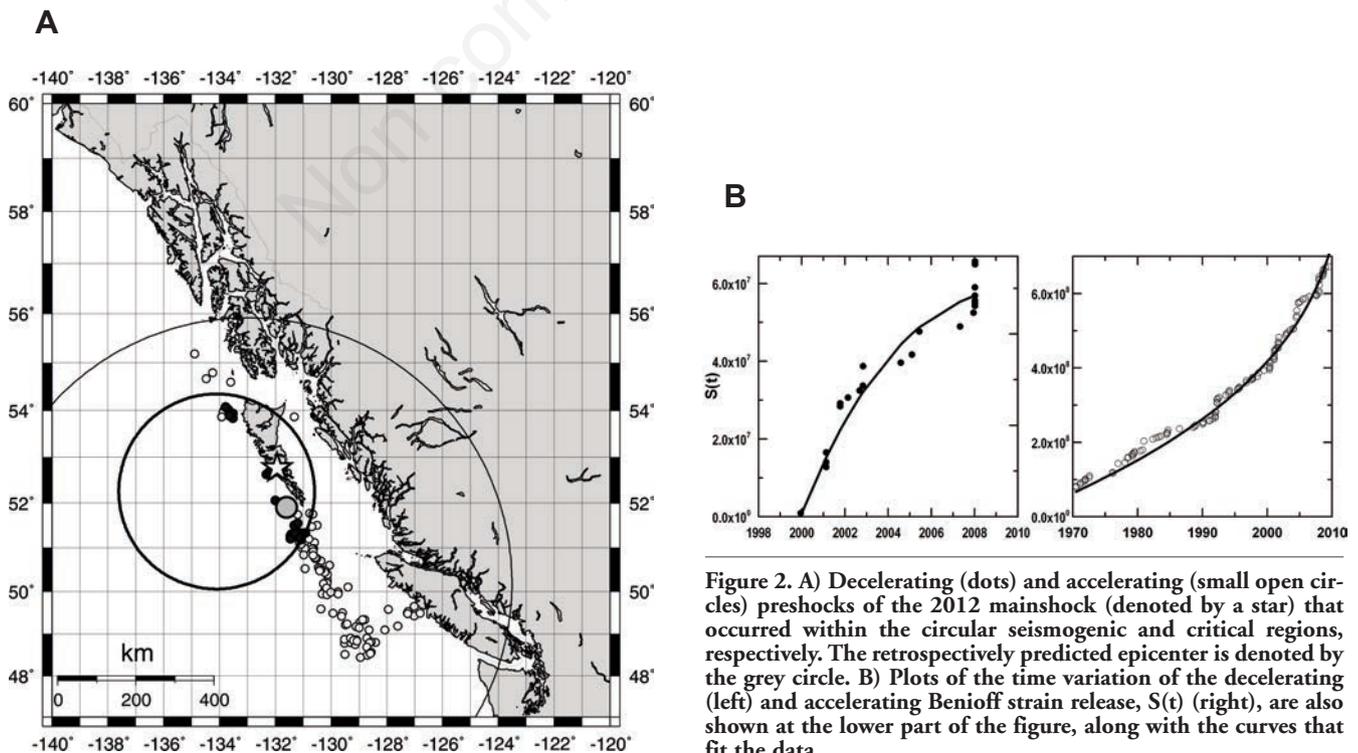


Figure 2. A) Decelerating (dots) and accelerating (small open circles) preshocks of the 2012 mainshock (denoted by a star) that occurred within the circular seismogenic and critical regions, respectively. The retrospectively predicted epicenter is denoted by the grey circle. B) Plots of the time variation of the decelerating (left) and accelerating Benioff strain release, $S(t)$ (right), are also shown at the lower part of the figure, along with the curves that fit the data.

$\varphi \pm 3^\circ$ NS and the meridians $\lambda \pm 3^\circ$ EW around the mainshock epicenter, $E(\varphi, \lambda)$. For this reason each of these areas is covered by a dense grid of points (e.g., $\pm 0.2^\circ$ NS, $\pm 0.2^\circ$ EW). In order to define the seismogenic region, where the decelerating Benioff strain release may be observed, each of these points is considered as the center of a circular seismogenic region with radius r (given by relation 2). For various r -values, t_{sd} start times and minimum magnitudes, different quality index, q_{di} , values are calculated. The geographic point where the largest of these values is found, corresponds to the center, F , of the seismogenic region. The same procedure is applied for the definition of the circular critical region with radius R (given by relation 3), and the geographic point where the largest quality index, q_{ci} , is calculated, corresponds to the center of the critical region where accelerating preshocks occurred prior to the mainshock generation.

Table 1 gives information on the properties of the decelerating and accelerating preshocks that preceded the 2012 Haida Gwaii mainshock and occurred within the corresponding seismogenic and circular regions. The first three columns list the origin time, magnitude and epicenter coordinates of the mainshock and the next two columns show the centers, F , Q , and radii r , R , of the circular seismogenic and critical regions, respectively. The minimum preshock magnitudes, M_{min} , and the start years of the decelerating, t_{sd} , and the accelerating, t_{sa} , preshock sequences are shown in the sixth

and seventh columns. The last three columns show the values of the curvature parameter, C , for the decelerating and accelerating preshocks, the logarithm of the strain rate, s_d and s_a (in Joules per year and per 10^4 km^2) and the values of the quality indexes q_a and q_d , given by the relations (6) and (7).

The spatial distribution of the decelerating (dots) and accelerating (small open circles) preshocks of the 2012 mainshock (denoted by a star) that occurred within the circular seismogenic and critical regions, respectively, is shown in Figure 2. The retrospectively predicted epicenter is denoted by the grey circle. Plots of the time variation of the decelerating (Figure 2B - left) and accelerating Benioff strain release, $S(t)$ (Figure 2B - right), are also shown at the lower part of the figure, along with the curves that fit the data. Both preshock sequences end three years before the mainshock.

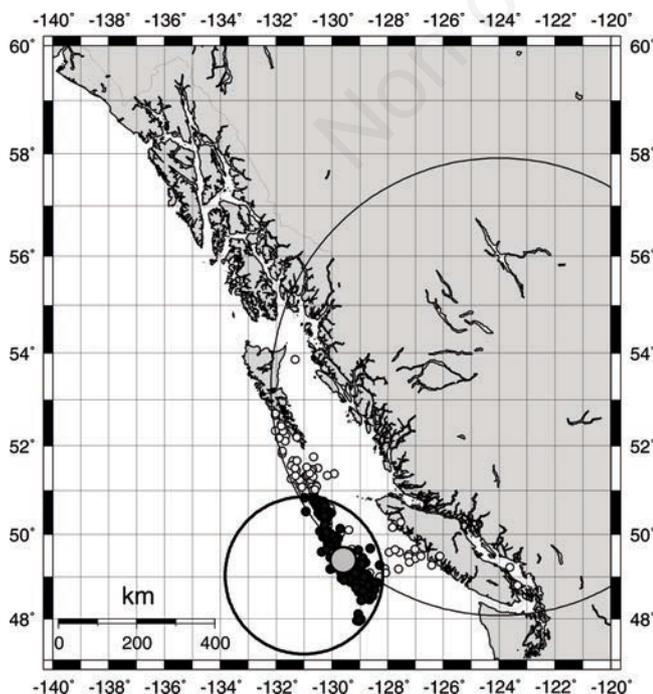
The finally adopted as predicted origin time, t^* , magnitude, M^* , and epicenter, $E^*(\varphi, \lambda)$, of the retrospectively predicted mainshock, which are given in the last row of Table 1, have been estimated in the following way. The origin time, t^* , is the mean value of the times calculated by the relations (4), (5) and (10) (the recurrence interval, T_r , calculated by the last relation, is added to the origin time of the previous mainshock that occurred in the region). The magnitude, M^* , is the mean of the magnitude values calculated by the relations (2), (3), and (11). Finally, the epicenter, $E^*(\varphi, \lambda)$, is the mean of two geographic points: i) the

mean epicenter of the decelerating preshocks; ii) the mean epicenter of the mainshocks which were identified after declustering the original complete catalogue of earthquakes that occurred in the seismogenic source. The difference between the observed and the calculated basic focal parameters, in the present work, are within the model uncertainties, i.e., $t^*: \pm 2.5$ years, $M^* \pm 0.3$, $E^*: 80 \pm 30 \text{ km}$.^{15,24}

In an attempt to search for precursory decelerating and accelerating preshock sequences that may be related to the generation of a future strong earthquake, the area bounded by the 48.0°N - 57°N parallels and 122°W - 138°W meridians was covered by a grid of points spaced 0.5° apart. The magnitude range of the probably ensuing earthquake was set between 7.0 and 8.2 with step equal to 0.2 magnitude units. Table 2 lists the basic parameters concerning decelerating and accelerating preshock sequences, which may be related to a probably ensuing mainshock. Figure 3 shows the spatial distribution of the decelerating and accelerating preshocks along with the time variation of the respective cumulative Benioff strain, $S(t)$.

Following the procedure described previously, the basic focal parameters of this earthquake were estimated: i) expected origin time $t_c^* = 2022.5$; ii) expected magnitude $M^* = 7.1$; iii) expected epicenter coordinates $E^* = 49.4^\circ\text{N}$, 129.6°W .

A



B

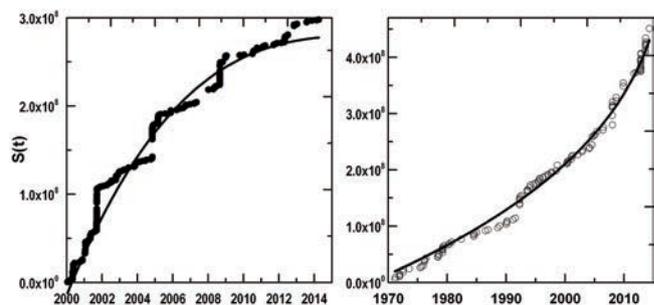


Figure 3. A) Decelerating (dots) and accelerating (small open circles) preshocks associated with the probably expected earthquake (grey circle) in area studied. B) Plots of the time variation of the decelerating (left) and accelerating Benioff strain release, $S(t)$ (right), are also shown at the lower part of the figure, along with the curves that fit the data.

Table 1. Properties of the decelerating and accelerating preshocks and their respective circular regions. Date, magnitude and epicenter coordinates of the 2012 mainshock are listed in the first three columns. The centers, F, and Q of the seismogenic and critical regions, along with their corresponding radii, r and R, are then shown. Information follows on the minimum magnitude, M_{\min} , the start times, t_{sd} and t_{sa} , and the values of the curvature parameter of the decelerating and accelerating preshocks. The last two columns show the logarithm of the strain rate, s_d and s_a (in Joules per year and per 10^4 km^2) and the values of the quality indexes q_d and q_a . In the last row the origin time, t^* , magnitude, M^* , and epicenter, $E^*(\varphi, \lambda)$, that were calculated in the present work, for this mainshock, are given.

Date	M	E (φ, λ)	F (φ, λ)	r (km)	M_{\min}	t_{sd}	C	$\log s_d$	q_d
2012.10.28	7.7	52.8-132.9	52.2-134.1	239	4.6	1998	0.53	5.19	3.6
			Q (φ, λ)	R (km)	M_{\min}	t_{sa}	C	$\log s_a$	q_a
			49.8-133.0	680	5.3	1968	0.38	5.18	7.3
t^*	M^*	$E^*(\varphi, \lambda)$							
2014.1	7.4	51.9-131.6							

Table 2. Properties of the decelerating and accelerating preshocks and their respective circular regions, which may be related to a probably ensuing strong earthquake in the area examined. In the last row the basic focal parameters $E^*(\varphi, \lambda)$, t^* , and M^* , of this earthquake are given.

F(φ, λ)	r (km)	M_{\min}	t_{sd}	C	$\log s_d$	q_d
52.2-134.1	239	4.6	1998	0.53	5.19	3.6
Q(φ, λ)	R (km)	M_{\min}	t_{sa}	C	$\log s_a$	q_a
49.8-133.0	680	5.3	1968	0.38	5.18	7.3
$E^*(\varphi, \lambda)$	t^*	M^*				
49.4-129.6	2022.5	7.1				

Discussion and Conclusions

We found that decelerating strain, released by intermediate magnitude preshocks, which occurred within a narrow region, along with accelerating strain released by larger shocks in a broader region, preceded the 2012 Haida Gwaii earthquake. We also found that the mainshocks, which occurred in the seismogenic source, where the 2012 epicenter is located, follow a quasi-periodic pattern which allows the estimation of occurrence time and magnitude of the next mainshock by the use of the magnitude and the occurrence time of the previous mainshock. Moreover, we searched the studied area for currently decelerating-accelerating strain release, which may be related with a probably ensuing mainshock.

We cannot ignore the skepticism expressed during the last several years about the hypothesis of accelerating seismicity that this pattern may arise from a combination of data fitting and from the generation of normal foreshocks and aftershocks⁹ or that it may be due to the formulation of the accelerating preshock generation as a power-law fit to a cumulative seismicity series.¹⁰ The negative results reported as regards the statistical significance of the AMR hypothesis,⁹ may be due to the optimization procedure applied, in examining acceler-

ating shocks that occurred in circular areas centered at the mainshock epicenter, which was based solely on minimizing the curvature parameter C. Recent tests on synthetic catalogues of earthquakes with spatiotemporal clustering based on the ETAS model in Aegean and California¹⁶ showed that decelerating and accelerating sequences are hardly recognizable in circular regions centered at the mainshock epicenter (low q_d and q_a values), in accordance with the negative results mentioned above,⁹ whereas such sequences are easily identifiable in circular areas with centers close to the mainshock epicenter.

References

- Cassidy JF, Rogers GC, Hyndman RD. An overview of the 28 October 2012 Mw7.7 earthquake in Haida Gwaii, Canada: a tsunamigenic thrust event along a predominantly strike-slip margin. *Pure Appl Geophys* 2014 [Epub ahead of print].
- Imamura A. *Theoretical and applied seismology*. Tokyo: Maruzen; 1937.
- Gutenberg B, Richter CF. *Seismicity of the Earth and associated phenomena*. New York: Hafner; 1954.
- Tocher D. *Seismic history of the San Francisco bay region*. Calif Div Mines Spec Rep 1959;57:39-48.
- Papadopoulos GA. Long-term accelerating foreshock activity may indicate the occurrence time of a strong shock in the Western Hellenic Arc. *Tectonophysics* 1988;152:179-92.
- Jaumé SC, Sykes LR. Evolving towards a critical point: a review of accelerating seismic moment/energy release rate prior to large and great earthquakes. *Pure appl Geophys* 1999;155:279-306.
- Papazachos CB, Karakaisis GF, Scordilis EM, et al. Global observational properties of the critical earthquake model. *B Seismol Soc Am* 2005;95:1841-55.
- De Santis A, Cianchini G, Di Giovabattista R. Accelerating moment release revisited: Examples of application to Italian seismic sequences. *Tectonophysics* 2015;639:82-98.
- Hardebeck JL, Felzer KR, Michael AJ. Improved tests reveal that the accelerating moment release hypothesis is statistically insignificant. *J Geophys Res* 2008;113: 808310.
- Mignan A. Retrospective on accelerating seismic release (ASR) hypothesis: Controversy and new horizons. *Tectonophysics* 2011;505:1-16.
- Bufe CG, Varnes DJ. Predictive modeling of

- seismic cycle of the Great San Francisco Bay Region. *J Geophys Res* 1993;98:9871-83.
12. Evison FF, Rhoades DA. The precursory earthquake swarm in New Zealand. *N Z J Geol Geophys* 1997;40:537-47.
 13. Evison FF. Long-range synoptic earthquake forecasting: an aim for the millennium. *Tectonophysics* 2001;333:207-15.
 14. Papazachos CB, Scordilis EM, Karakaisis GF, et al. Decelerating preshock seismic deformation in fault regions during critical periods. *B Geol Soc Greece* 2005;36:1490-8.
 15. Papazachos CB, Karakaisis GF, Scordilis EM, et al. New observational information on the precursory accelerating and decelerating strain energy release. *Tectonophysics* 2006;423:83-96.
 16. Karakaisis GF, Papazachos CB, Scordilis EM. Recent reliable observations and improved tests on synthetic catalogs with spatiotemporal clustering verify precursory decelerating-accelerating seismicity. *J Seismol* 2013;17:1063-72.
 17. Papazachos BC, Karakaisis GF, Papazachos CB, et al. Evaluation of the results for an intermediate term prediction of the 8 January 2006 Mw=6.9 Cythera earthquake in southwestern Aegean. *B Seismol Soc Am* 2007;97:347-52.
 18. Papazachos BC, Karakaisis GF, Papazachos CB, et al. A forward test of the decelerating-accelerating seismic strain model in the Mediterranean. *Boll Geofis Teor Appl* 2009;50:235-54.
 19. Bowman DD, Quillon G, Sammis CG, et al. An observational test of the critical earthquake concept. *J Geophys Res* 1998; 103:24359-72.
 20. Papazachos CB, Karakaisis GF, Savvaidis AS, et al. Accelerating seismic crustal deformation in the southern Aegean area. *B Seismol Soc Am* 2002;92:570-80.
 21. Papazachos BC, Papadimitriou EE, Karakaisis GF, et al. Long-term earthquake prediction in the Circum-Pacific convergent belt. *Pure Appl Geophys* 1997;149:173-217.
 22. Papazachos BC, Papaioannou ChA. Long term earthquake prediction in the Aegean area based on the time and magnitude predictable model. *Pure Appl Geophys* 1993;140:593-612.
 23. Kagan YY, Jackson DD. Long-term earthquake clustering. *Geophys J Int* 1991;104:117-33.
 24. Papazachos BC, Karakaisis GF, Scordilis EM. Time dependent seismicity in the continental fracture system. *Boll. Geofis Teor Applic* 2014;55:617-39.
 25. Halchuk S. Seismic Hazard Earthquake Epicenter File (SHEEF) used in the fourth generation seismic hazard maps of Canada. Geological Survey of Canada, Open File 6208 2009; pp 16.
 26. Milne WG, Rogers GC, Riddihough RP, et al. Seismicity of western Canada. *Can J Earth Sci* 1978;15:1170-93.
 27. Wiemer S. A software package to analyze seismicity: ZMAP. *Seis Res Letts* 2001;72: 374-83.
 28. Gutenberg B, Richter CF. Frequency of earthquakes in California. *B Seismol Soc Am* 1944;34:185-8.
 29. Scordilis EM. Globally valid relations converting Ms, mb and MJMA to Mw, at Meeting on earthquake monitoring and seismic hazard mitigation in Balkan countries. NATO ARW, 11-17 September 2005, Borovetz, Bulgaria, pp 158-161.
 30. Scordilis EM. Empirical global relations converting Ms and mb to moment magnitude. *J Seismology* 2006;10:225-36.
 31. Ristau JP. Seismotectonics of western Canada from regional moment tensor analysis. PhD Thesis, University of Victoria, Victoria, Canada; 2004, 209 pp.
 32. Papazachos BC, Papaioannou ChA, Scordilis EM, et al. A forward test of the decelerating-accelerating seismic strain model to western south and Central America. *Tectonophysics* 2008;454:36-43.